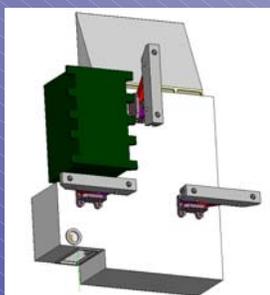


How Visible is Regional Smog Ozone from Space and How Can We Make It More Visible?

Robert Chatfield, NASA Ames

Robert Esswein, NASA Ames and BAER

Mark Schoeberl, NASA GSFC

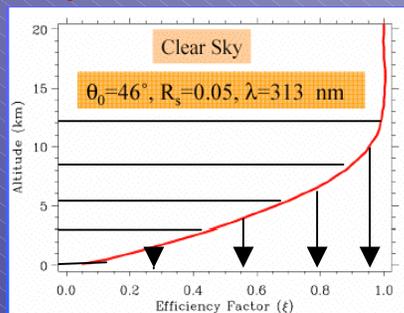


(1) Retrieval theory says that middle- and upper-tropospheric air should dominate

- *What do sonde comparisons tell us about vertical response of the tropospheric ozone estimates?*

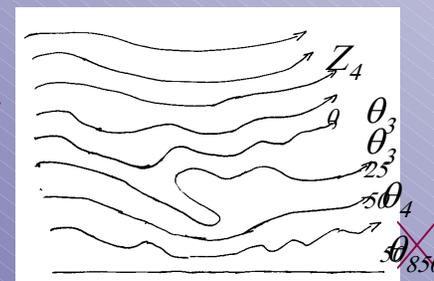
(2) We propose an ozone using AIRS depictions of stratospheric structure (T , Z , θ) and correlations/regressions, not (yet) AIRS O_3 . Conceivably, this could remove low tropopause situations and even UT fronts.

- *How well do this estimate work (so far) and what is its vertical response function, based on IONS-06 sondes?*



Strat.

Trop.



Aura Science Team Meeting
Pasadena, October 2, 2007



Two Concerns We Address:

(1) Retrieval theory says that middle- and upper-tropospheric air should dominate tropospheric column retrievals

- What do sonde comparisons tell us about vertical response of the tropospheric ozone estimates?

(2) We propose an ozone using AIRS depictions of stratospheric structure (T , Z , θ) and correlations/regressions, not (yet) AIRS O_3 . Conceivably, this could remove low tropopause situations and even UT fronts.

- How well do this estimate work (so far) and what is its vertical response function, based on IONS-06 sondes?

In South, OMI-MLS informs about smog, Where about tropopause and fronts

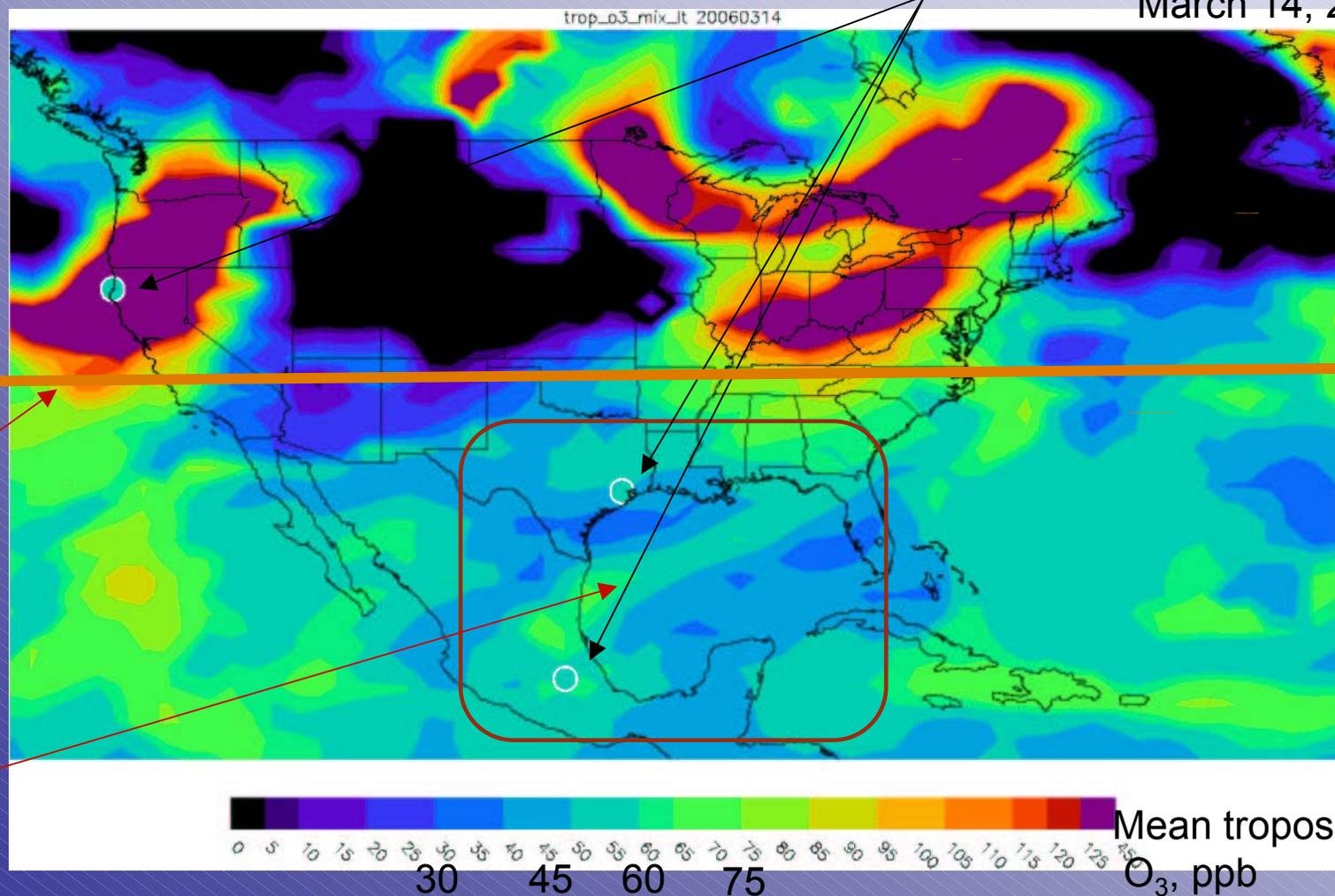
OMI-MLS tropospheric O₃ as estimated for a INTEX-B Mexico City plume event. Above the orange line, the sensitivity to ozone below ~200 hPa commonly produces very high ozone features which race W to E. South of the line, features move more slowly, such as the Central-Mexico plume.

Schoeberl Tropospheric Ozone Column to estimated tropopause
Schoeberl pre-Mar07 version
Subtropics easier to interpret, compare

IONS Sondes comparisons

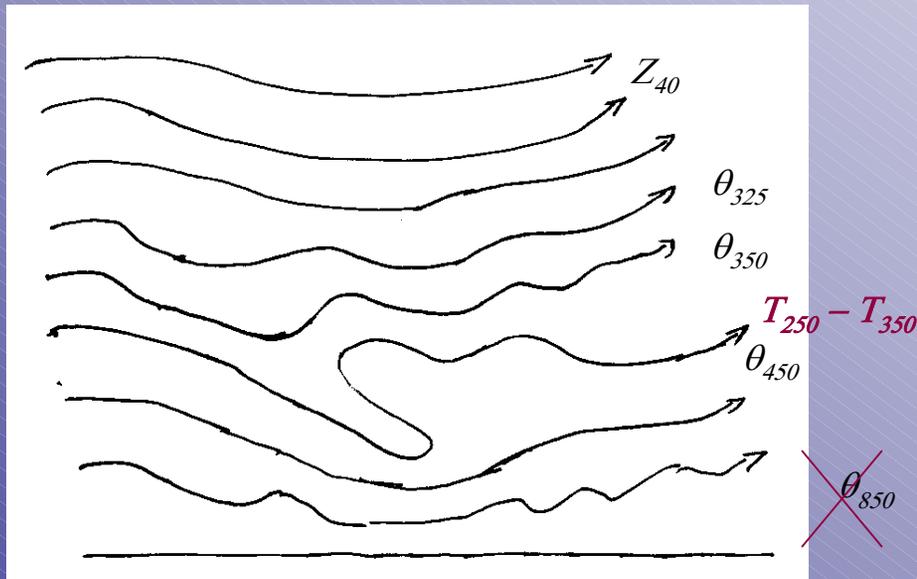
(best when colors match)

March 14, 2006



OMI-AIRS(θ):

Tropospheric Ozone from OMI Total Column O₃ and AIRS Temperatures : Stratosphere and “Distraction” Removal:



Specific technique “projection pursuit regression” – roughly analogous to empirical orthogonal functions, but assembling combinations of those explaining variables (θ 's) that have the most explanatory value for the explained variable, total ozone.

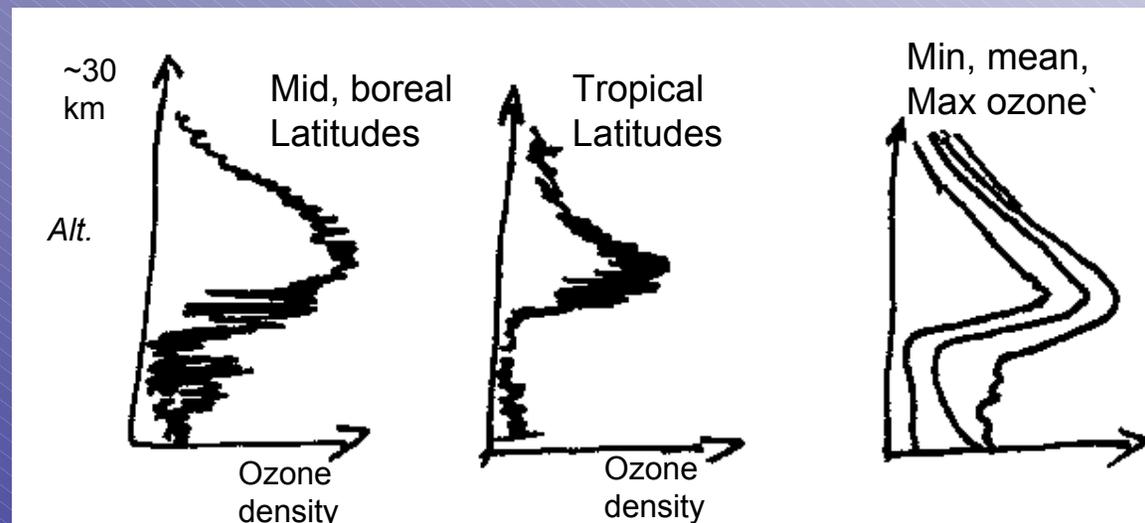
Sums: average properties of layers

Differences of T 's: Lapse rate

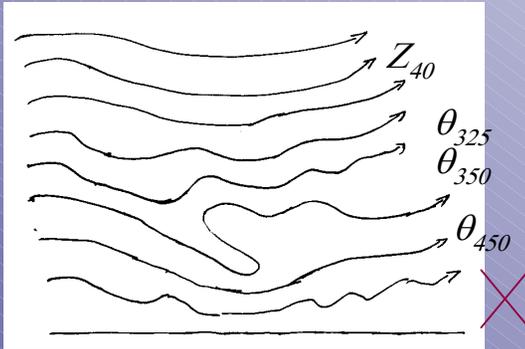
Do not include lower-tropos. variables

Some supporting rationalization is that over periods of a month or so, O₃ - θ relationships characterize the stratosphere vertically and regionally; also that local positions of the tropopause and UT fronts are *fairly well* captured by AIRS.

T 's, θ 's, Z 's (all are varied expressions of AIRS temperature structure.)



Example of Stratospheric Fit: August, 2006

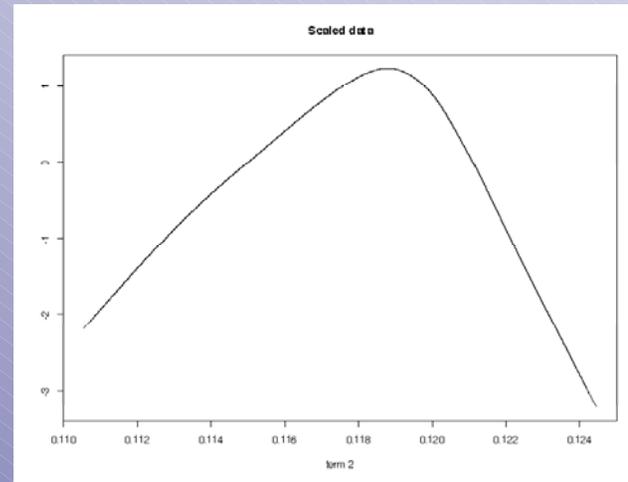
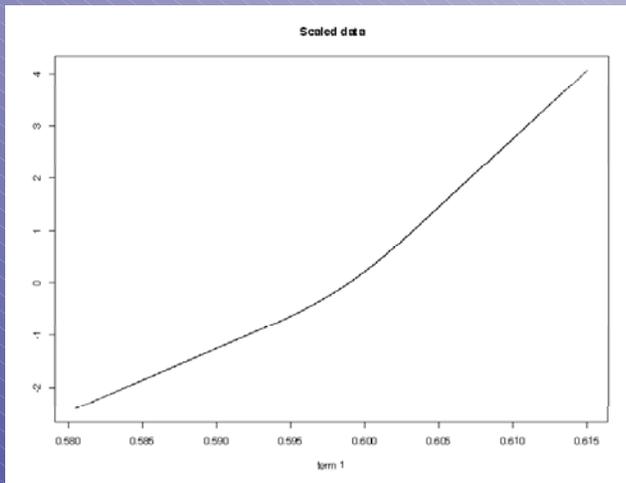


T 's on a p -surface map directly to θ 's; Z 's are simply integrals of the T information.
Vertical differences in T define tropopause and front:
Limitation: limited AIRS vertical resolution for T .

AIRS Variable	term 1	term 2	term 3	term 4	term 5	term 6
T_{450}	-0.01	-0.03	-0.01	-0.20	-0.07	0.33
T_{350}	-0.33	-0.02	0.04	0.03	0.40	-0.01
T_{275}	0.25	0.06	-0.01	-0.15	-0.59	0.40
T_{225}	-0.22	-0.06	0.04	-0.04	0.29	-0.20
Z_{175}	-0.23	-0.03	0.01	0.34	0.18	-0.53
Z_{60}	0.45	0.79	-0.73	-0.58	-0.26	0.58
Z_{40}	0.72	-0.61	0.68	0.70	-0.55	-0.28
latitude	-0.02	0.01	0.00	-0.01	0.01	0.01
Relative Contribution	0.05	0.03	0.02	0.01	0.01	.004
249317 fit points	$r = 0.92$	$r^2 = 0.84$				

- Why such a large variance explained? => Both vertical and horizontal (N/S) fit
- Any fitting of the troposphere is accidental ... accidents can happen: I.e., correlations of structure between troposphere and stratosphere (particularly UT?)

Example of Stratospheric Fit: First and second combinations



Response function for linear combinations of variables. In this situation, where we expect stratospheric ozone and stratospheric dynamical variables to exhibit simple relationships, the response curves are simple combinations of linear functions: recall that the first projection pursuit directions attempt to summarize all the most significant information, and often these are simple combinations or contrasts (differences)

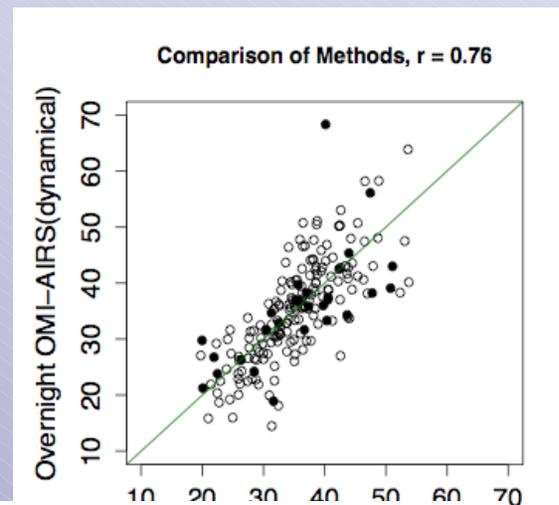
PPR also allows us to organize somewhat correlated information from several levels into an informative few variable sums

PPR gives “semi-interpretable fits” suggesting more analysis: Some other modern methods may speed interpretation.

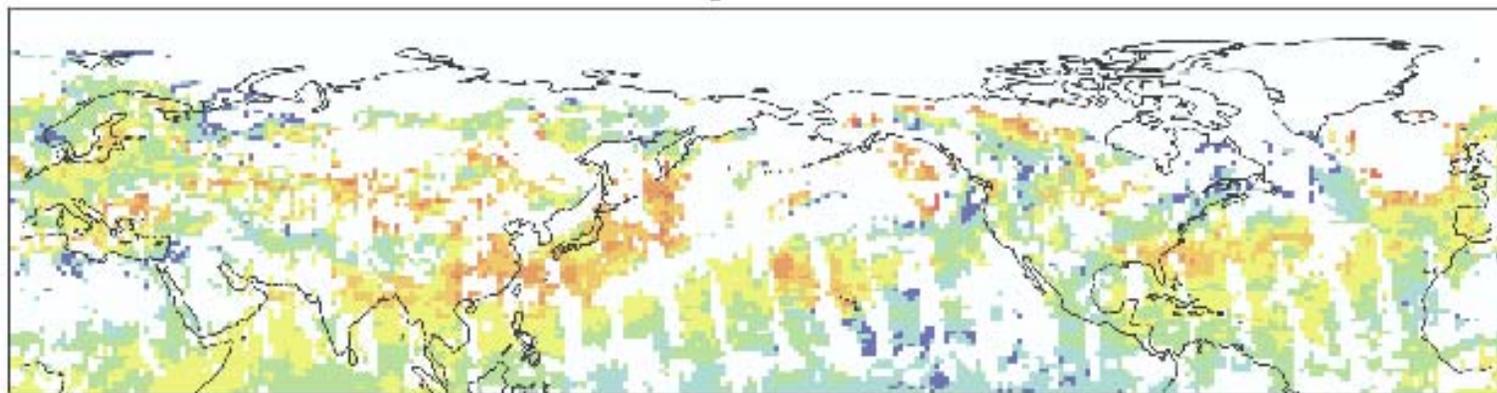
Comparison of the “tropospheric” OMI-MLS-trajectory and the OMI-AIRS(θ) are as good as expectable

However, not all high values (UT-Strat features?) are removed.
Comparisons sem better than comparisons of either with sondes!

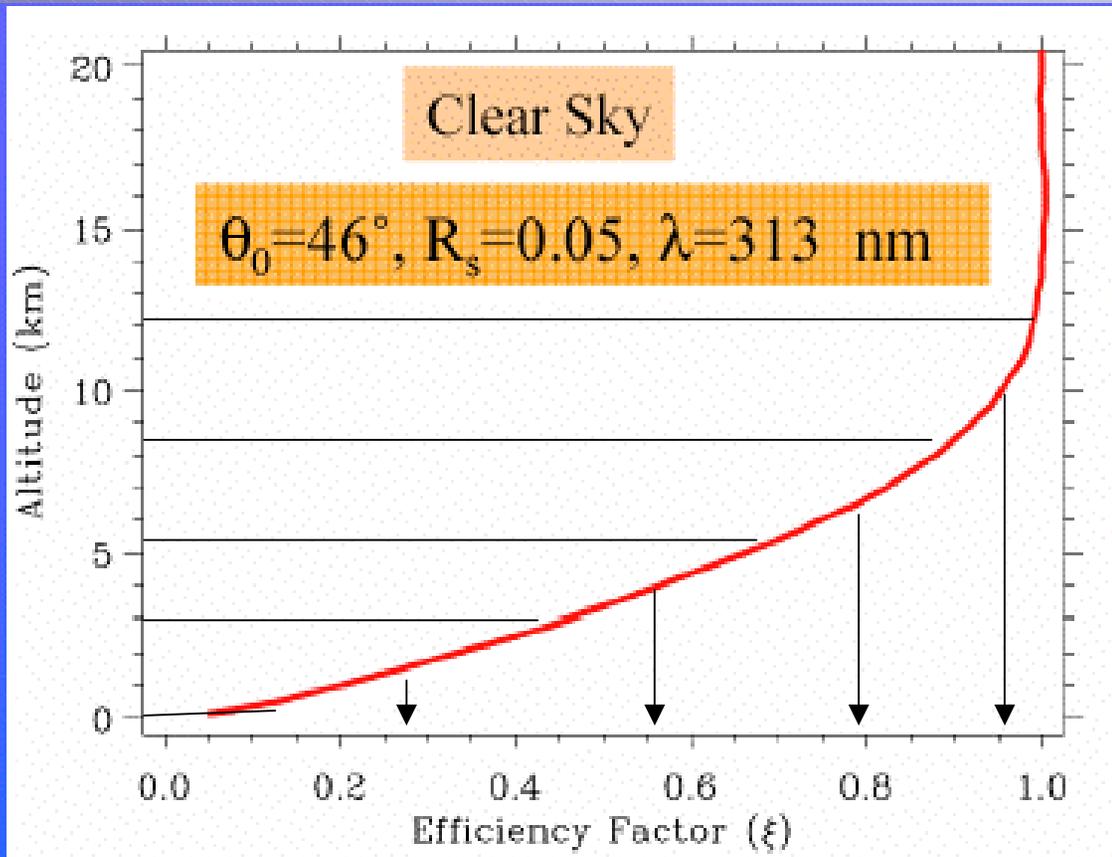
Clouds reported by OMI and AIRS must be understood and rationalized, so that regions of the troposphere predominantly clear over 4 km can be



O3 DU Corrected Avg 20060504 and 20060505



Theoretical clear-sky response of OMI total ozone column to lower-tropospheric ozone is low!



Theoretical sensitivity of the UV-based total ozone measurement to tropospheric ozone, decreasing greatly below 5-6 km. Reproduced from PK Bhartia, this sensitivity includes known basic scattering-absorption physics for the situation shown.

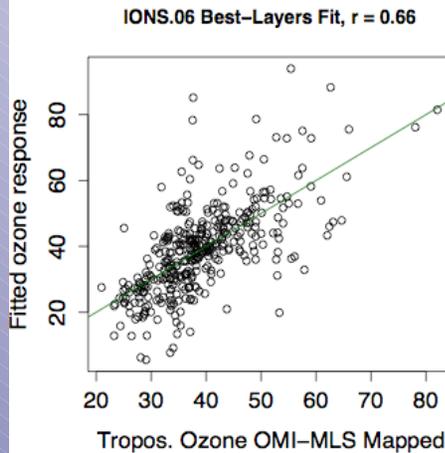
Our analysis model layers for sonde comparisons are shown: for the Sfc-700 hPa region 1 DU true ozone (e.g., estimated by sondes) should produce 0.3 DU total column ozone.

Strong covariation of 0-3 km (Sfc-700 hPa) ozone with OMI-based estimates

Comparison with layer-averaged ION2-2006 Sondes

Multiple regression analysis of OMI-based tropospheric column ozone with layer column averages gives near-1:1 contribution in lower troposphere, similar or high in UT, lower contribution in mid-tropospheric layers.

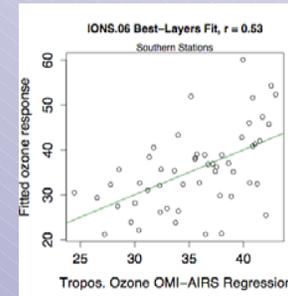
Why? Are there statistical effects connected with aerosol scattering, cloud-top scattering, or treatments of cloudiness?



OMI-MLS (Schoeberl Estimates)

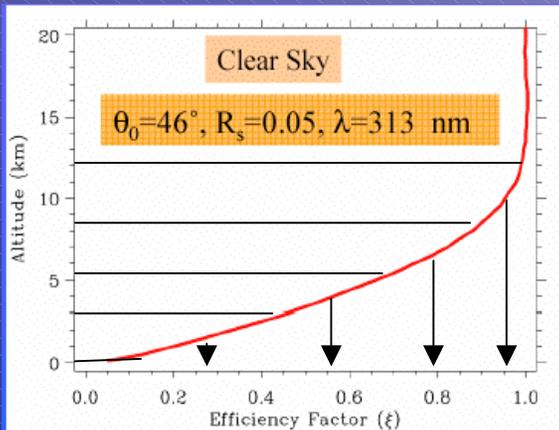
Beltville, Boulder, Brattslake, Edmonton, Egbert, Kelowna, Narragansett, Native, Paradox, Sable, Trinidad, Valparaiso, Wallops, Walsingham, Yarmouth

similar correlation, no very high values



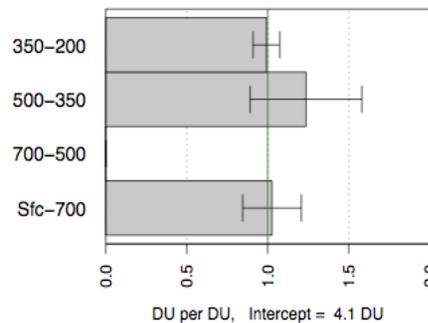
OMI-AIRS(θ)

August only
NORTHERN STATIONS



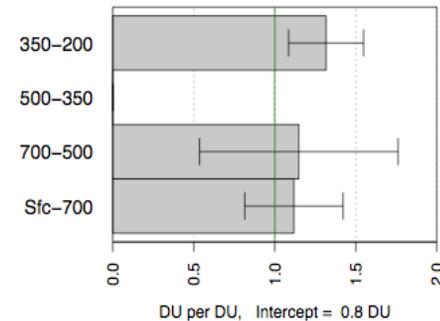
IONS.06 Reg. Coeff. and Std. Err.

Northern Stations, OMI-MLS



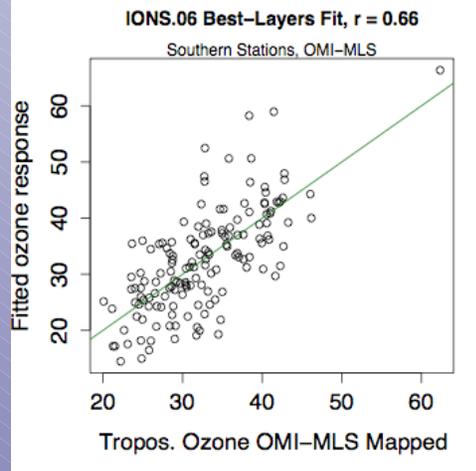
IONS.06 Reg. Coeff. and Std. Err.

Northern Stations

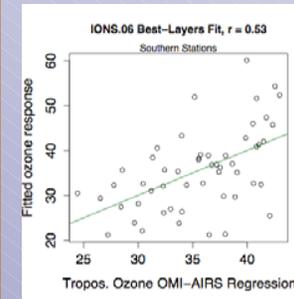


Southern IONS Ozone Soundings: Strong covariation of 0-3 km (Sfc-700 hPa) ozone with OMI-based estimate

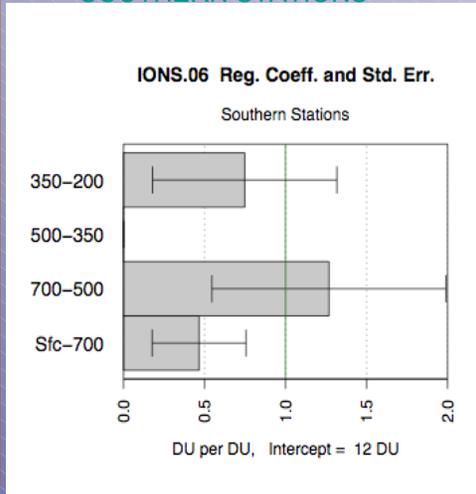
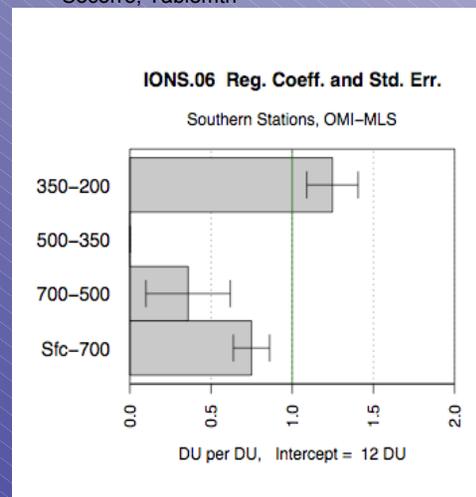
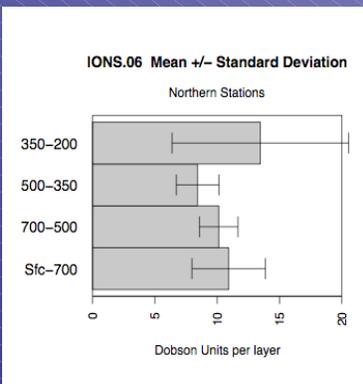
- Best-fit regressions are shown (chosen using Mallows Cp statistic).
- Note low intercepts
- New method: August IONS (TexAQ5-2006) only: more months in progress.
- Means DU in each layer very similar



OMI-MLS (Schoeberl Estimates)
SOUTHERN STATIONS:
Holtville, Houston, Huntsville, Mexico,
Socorro, Tablemtn



OMI-AIRS(θ)
August only
SOUTHERN STATIONS



End

Extra slides follow

Conclusions:

“Best-regression” studies suggest that O₃ in the lower 3 km of the troposphere contributes significantly both to **OMI–MLS (mapped) column estimates** and **a new method** often with near 1:1 weighting. Air in the 350-200 hPa region contributes similarly, However, *mid-tropospheric ozone often contributes weakly, insignificantly, occasionally negatively (in estimates)*. Other regression methods agree.

A tropospheric-ozone estimate using **AIRS depictions of evolving stratospheric dynamical structure (T, Z, θ)** and correlations/regressions compares well to sonde and MLS-based estimates. Some UT/LS high-ozone remains, but is reduced

- **Considerable improvement is possible** to this variance-based technique.

GEO Performance Data

Measurements: -- □field of regard = 22° diameter and
 □footprint size *at nadir* = **2.5 km @ 2.3 μm**;
5.0 km @ 3.6 and 4.7 μm;
 and **10.0 km @ 9.6 μm**

□Areal coverage = 2500 km x 2500 km per 20 minutes

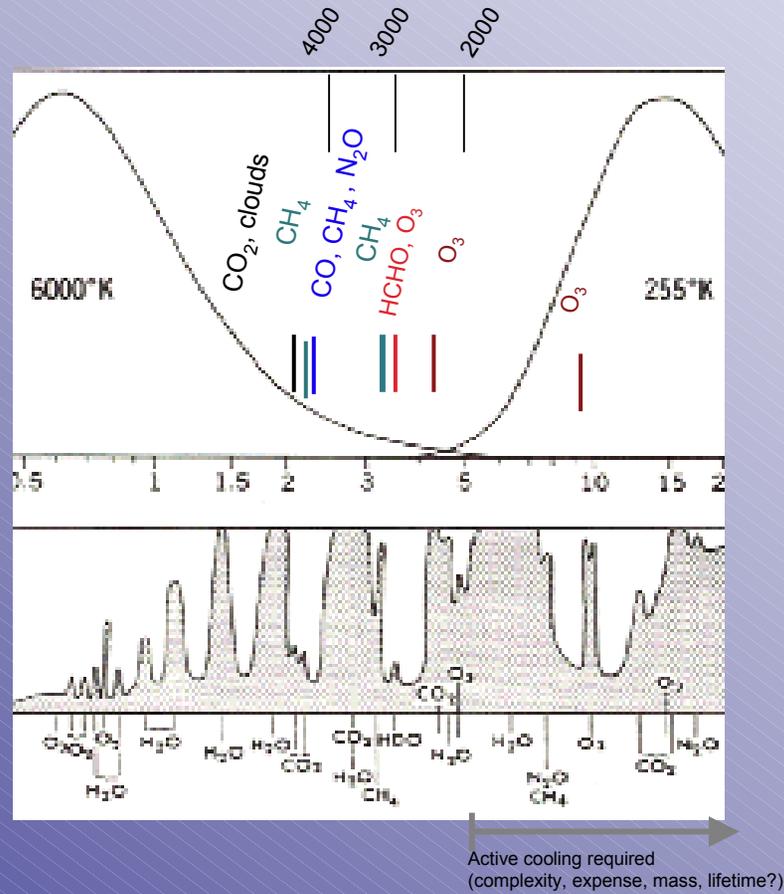
□ Threshold spectral range $\nu_1 \rightarrow \nu_2$, resolution ($\Delta\nu$) & NEdN characteristics

channel	$\nu_1 \rightarrow \nu_2$ (cm ⁻¹)	$\Delta\nu$ (cm ⁻¹)	NEdN (nW/(cm ² sr cm ⁻¹))
~ 2.3 μm	4281 to 4301	0.13	1.0
~ 3.6 μm	2778 to 2791	0.13	1.0
~ 4.7 μm	2112 to 2160	0.20	1.0
~ 9.5 μm	1043 to 1075	0.10	2.0

Retrieval expectations:

- **O₃ including the BL** and 3 additional layers below 22 km with precision <5% in the latter:
 i.e., < 5% in 0 to ~3 km region: both column (3.6) and thermal (9.5) contribute
 Aggregating four 5-km footprints should
- **CO in the BL** and 2 layers above with respective precisions the order 10, 5 and 3%
- **HCHO** with column precision < 4 x 10¹⁵ /cm².
 Height information from day and night retrievals.
- Some CH₄ information should be available: this was a pollutio-oriented

Wavelengths and Species (ii)

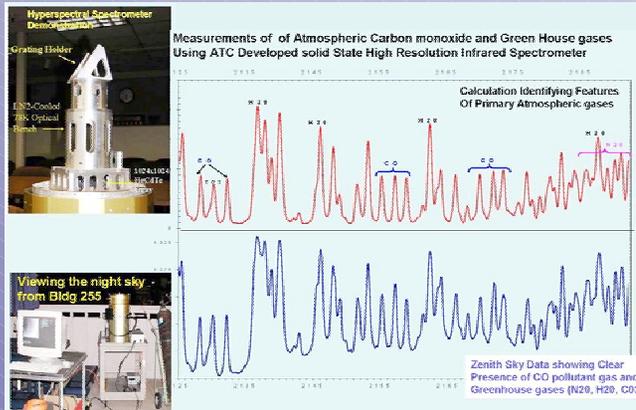


HQ Costing Study: Alternative Instrument Concepts (2/2)

GEO TIMS

Mass: est 87 kg
Power: est 160 W
Volume: 0.43 m x
 0.24 m x 0.67m

TIMS = Tropospheric
 Infrared mapping
 Spectrometers



Clear sky spectral data near 4.7 um compared with a model. Data, with $\nu \sim 0.5 \text{ cm}^{-1}$, were obtained with demonstration predecessor to the IIP GMS.

Features of the TIMS Measurement Concept

- Employs two grating mapping spectrometers (GMS);
- Utilize separate 9 cm apertures & scan mirrors
 - Each has 2 channels: 3.6 & 2.3 μm and 9.6 & 4.7 μm
 - 3.6 μm channel uses solar reflective (SR) and thermal IR signal to obtain
 - Column O_3 with sensitivity in the BL
 - **HCHO** in full and partial column
 - The 9.6 μm channel provides layers of O_3 in the troposphere and above, and also
 - BL O_3 by combined retrieval with SR data
 - The 2.3 and 4.7 μm channels provide **CO** in 3 layers, including the BL with precision better than 10%.
 - Ancillary retrievals of BL & profile **CH₄** and **H₂O**, and **N₂O** & **CO₂** column

Performance Data

- Measurements:** -- □ field of regard = 22° diameter and □ footprint size @ nadir =
 2.5 km @ 2.3 μm ; 5.0 km @ 3.6 and 4.7 μm ; and 10.0 km @ 9.6 μm
- Areal coverage = 2500 km x 2500 km per 20 minutes
 - Threshold spectral range $\nu_1 \rightarrow \nu_2$, resolution ($\Delta\nu$) & NEΔN characteristics

channel	$\nu_1 \rightarrow \nu_2$ (cm^{-1})	$\Delta\nu$ (cm^{-1})	NEΔN ($\text{nW}/(\text{cm}^2 \text{sr cm}^{-1})$)
~ 2.3 μm	4281 to 4301	0.13	1.0
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~ 9.5 μm	1043 to 1075	0.10	2.0

Retrieval expectations:

- **O₃** including the BL and 3 additional layers below 22 km with precision <5% in the latter
- **CO** in the BL and 2 layers above with respective precisions the order 10, 5 and 3%
- **HCHO** with column precision < $4 \times 10^{15} / \text{cm}^2$.

Technology Development Needs

1. IIP demonstration (2006-2008) of the TIMS GMS will result in TRL 5+
 - a. Includes GMS operating near 2.3 μm & 4.7 μm .
 - portable to facilitate field measurements
 - b. CO retrieval from atmospheric measurements
 - validated by retrievals with data from Denver University FTS
2. 9.6 μm channel demonstration
 - a. Large format, low noise array with cutoff ~ 10.5 μm
 - b. Suitable detector array has been demonstrated on a high noise direct injection mux
 - we anticipate no problem on low noise, low light mux

➤ **Not chosen for design study owing to relatively lower TRL, ~4, but should be considered for future.**